

1 **Spatiotemporal variability of submicrometer particle number**
2 **size distributions in an air quality management district**

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2 **Abstract**

3 First measurements of ambient 10-1000 nm particle number concentrations (N_{TOT})
4 and size distributions were made at an urban, coastal, mountain and downwind site within
5 the Central Taiwan Air Quality Management District during a cold and a warm period.
6 The primary objectives were to characterize the spatial and temporal variability of the
7 size-fractionated submicrometer particles and their relationships with copollutants and
8 meteorological parameters. The results show that the ultrafine particles (<100 nm) are the
9 major contributor to the N_{TOT} . The mean N_{TOT} was highest at the urban site, whereas
10 lower and comparable at the three other sites. Although the mean N_{TOT} at each site
11 showed insignificant differences between study periods, their diurnal patterns and size
12 distribution modal characteristics were modestly to substantially different between study
13 sites. Correlation analyses of time-resolved collocated aerosol, copollutants and
14 meteorological data suggest that the observed variability is largely attributable to the
15 local traffic and to a lesser extent photochemistry and SO_2 possibly from combustion
16 sources or regional transport. Despite sharing a common traffic source, the ultrafine
17 particles were poorly correlated with the accumulation particles (100-1000 nm), between
18 which the latter showed strong positive correlation with the $PM_{2.5}$ and PM_{10} . Overall, the
19 N_{TOT} and size distributions show modest spatial heterogeneity and strong diurnal
20 variability. In addition, the ultrafine particles have variable sources or meteorology-
21 dependent formation processes within the study area. The results imply that single-site
22 measurements of $PM_{2.5}$, PM_{10} or N_{TOT} alone and without discriminating particle sizes
23 would be inadequate for exposure and impact assessment of submicrometer particle
24 numbers in a region of diverse environments.

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26 Keywords: nanoparticles, number size distribution, spatial distribution, temporal
27 variability, exposure assessment, emission sources

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1 **1. Introduction**

2 Atmospheric particles have significant implications in climate, environmental and
3 human health. These microscopic airborne particles, on a global scale, influence climate
4 by absorbing and scattering solar radiation, modifying cloud formation, lifetime and
5 precipitation (Rosenfeld et al., 2008). On a regional scale, they degrade air quality and
6 impair visibility (US EPA, 2010a). On a microscopic scale, elevated particle
7 concentrations play a major role in exerting adverse human health effects (US EPA,
8 2010b). The extent of those effects depends on the particle size, concentration and
9 chemical composition (e.g., Dusek et al., 2006; Zanobetti and Schwartz, 2009; Peng et al.,
10 2009). Studies have shown smaller particles are more pertinent to adverse human health
11 effects, thereby resulting in a worldwide regulatory focus shift from TSP to PM₁₀, and
12 then to PM_{2.5}. More recently, ultrafine particles (UFP, < 100 nm) are attracting growing
13 concerns due to their potential higher toxicity and ability to deposit efficiently in the
14 alveolar region upon inhalation (Nel et al., 2006).

15 Numerous studies have implicated primary vehicle emissions and secondary new
16 particle formation (NPF) as the two major sources of UFP. A comprehensive review paper
17 of ambient UFP from the two sources has been given by Morawska et al. (2008). In
18 addition to sources, meteorological and environmental factors such as wind speed,

1 temperature, relative humidity, mixing height, local-scale roadside structures, interactions
2 of multiple streets and topography also could play an important role in modulating the
3 spatial and temporal variability of particle number size distributions (Zhu et al., 2002;
4 Wehner and Wiedensohler, 2003; Zhu et al., 2004; Paatero et al., 2005; Hussein et al.,
5 2006; Bowker et al., 2007; Ogulei et al., 2007; Jamriska et al., 2008; Wang et al., 2011a;
6 Wang et al., 2011b; Sabaliauskas et al., 2011).

7 Strong temporal variability of particle number has been widely reported, especially
8 for the UFP as compared to the accumulation particles (e.g., Wehner and Wiedensohler,
9 2003; Wang et al., 2011b). In urban environments, the diurnal pattern of particle number
10 is typically characterized by a peak concentration in the morning and another in the early
11 or late evening (Jeong et al., 2004; Park et al., 2008; Lonati et al., 2011; Wang et al.,
12 2011b). The former is particular distinct in traffic-impacted areas, whereas the latter is
13 less consistent due to the variable sources and meteorological conditions in different
14 areas. In addition, a third peak concentration due to NPF event could occur around
15 midday or in the early afternoon. This is the time of day when photochemistry is most
16 intense and the increase of mixing height lowers the pre-existing particle concentration
17 (Young and Keeler, 2007; Moore et al., 2007; Park et al., 2008; Wang et al., 2010; Wang
18 et al., 2011b). The threshold of pre-existing particle concentration for NPF is not an

1 absolute value but a relative term, which additionally needs to take into account of the
2 condensable vapor concentration (e.g., H₂SO₄) (Stanier et al., 2004a; McMurry et al.,
3 2005). For example, Kulmala et al. (2004) have reported that NPF events were observed
4 in all types of environment, i.e., not limited to clean environments but included polluted
5 ones (Zhu et al., 2004; Dunn et al., 2004; Monkkonen et al., 2005; Wu et al., 2007; Yao et
6 al., 2010). On a seasonal basis, a number of studies have shown that the particle number
7 concentrations are lowest in the summer and highest in the winter. In addition, the
8 smallest size fraction (<50 nm) of UFP shows the strongest seasonal fluctuations (Wehner
9 and Wiedensohler, 2003; Jeong et al., 2004; Harrison and Jones, 2005; Wang et al., 2010;
10 Wang et al., 2011b). Poor atmospheric mixing and low temperatures are the two main
11 reasons leading to elevated particle number concentration in the winter.

12 Spatial variability of particle number concentrations has been commonly studied by
13 taking measurements as a function of distance from the emission source. These studies
14 have shown a continuous decrease of particle number concentrations while the particles
15 advect away from the road (e.g., Zhu et al., 2002; Zhu et al., 2004; Kittelson et al., 2004;
16 Westerdahl et al., 2005; Buonanno et al., 2009). In a downwind buffer zone of no traffic
17 and little obstruction from buildings, Zhu et al. (2002; 2004) showed that the roadway
18 particle number, CO and black carbon decrease exponentially with distance and then

1 become indistinguishable from the background level at a distance of 300 m. Furthermore,
2 the decrease rate is slower in the winter than in the summer. The resulting changes in the
3 number size distribution during the road-to-ambient process were attributed largely to
4 condensation, evaporation and dilution, and to a lesser extent to coagulation and
5 deposition (Zhang et al., 2004). In a more complex environment of multiple streets,
6 Ogulei et al. (2007) observed much different patterns of source impacts on the particle
7 number size distributions. Using highly time-resolved data, the authors were able to show
8 distinction between fresh versus aged diesel particles and spark- versus compress-ignition
9 engine emissions. Another way to determine spatial variability is to take measurements at
10 several locations within a city or region. However, data from this type of study are still
11 relatively sparse. Harrison and Jones (2005) showed that the particle number
12 concentrations at seven out of eight sites in the UK were generally similar, suggesting
13 little inter-urban variability in the urban air. One-year concurrent measurements of
14 particle number at five different sites in the Los Angeles Basin showed low to modest
15 spatial correlations for particle number (Sardar et al., 2004). Moore et al. (2009) found
16 the intra-community variability of UFP concentrations at 14 sites in urban Los Angeles is
17 comparable to and exceeds the inter-community variability presented in earlier studies. In
18 the same study, the concentration variability of smaller particles (< 40 nm) was found to

1 be higher than larger particles (Krudysz et al., 2009). A follow-up study by Hudda et al.
2 (2010), with additional measurements at receptor sites up to 115 km downwind to Los
3 Angeles, also suggested that the intra-community variability was larger than the inter-
4 community variability. More recently, Wang et al. (2011b) showed that, despite uniform
5 temporal variation, the UFP present moderate spatial divergence in urban Rochester, New
6 York. The abovementioned spatial variability has been the largest uncertainty in exposure
7 estimates for population-based epidemiological studies (US EPA, 2010b).

8 Earlier studies of UFP in Taiwan have focused primarily on particle mass and
9 chemical composition (e.g., Lin et al., 2005; Hsieh et al., 2009; Chen et al., 2010a).
10 Measurements of ambient submicrometer or UFP number size distributions are lacking
11 until recent years (Chang and Lee, 2007; Chen et al., 2010b; Cheng et al., 2010a; Cheng
12 et al., 2010b; Chen et al., 2011). Despite these progresses, the spatial and temporal
13 variability of particle number size distributions has yet been addressed. With that in mind,
14 the present study selected four sampling sites in the Central Taiwan Air Quality
15 Management District (CTAQMD) for characterizing the spatiotemporal variability of
16 submicrometer particle number size distributions and their relationships with copollutants
17 and meteorological parameters. The selected sites encompass relatively diverse
18 environments, including an urban, coastal, mountain and downwind area. Measurements

1 were carried out consecutively among the sites during a cold and a warm period. In
2 addition to complementing earlier studies, the results of the present study will serve as the
3 foundation for future in-depth particle number size distribution studies. The outputs are
4 pertinent to obtaining a more representative assessment of their impact on human and
5 environmental health.

6 **2. Material and Methods**

7 **2.1 Sampling Sites**

8 The aerosol sampling campaigns took place in the CTAQMD, one of the seven air
9 quality management districts designated by the Taiwan Environmental Protection
10 Administration (TW EPA) (<http://www.epa.gov.tw/>). The CTAQMD includes the Greater
11 Taichung City, Nantou County and Changhua County, and covers an area of $\sim 7,400 \text{ km}^2$
12 with a registered population of $\sim 4,480,000$. It is to be noted that the registered number of
13 vehicles is approximately the same as the registered population. The highest population
14 density of $\sim 6570 \text{ inhabitant km}^{-2}$ is in the Taichung urban area, whereas the lowest of \sim
15 $130 \text{ inhabitant km}^{-2}$ in Nantou County.

16 Four out of the 12 TW EPA air quality monitoring sites in the CTAQMD were
17 selected to nominally represent urban, coastal, mountain, and downwind areas. They

1 include Jhongming (JM), Siansi (SS), Puli (PL), and Jushan (JS) site, respectively. The
2 sites are 20-50 km apart as shown in Figure 1 and the site-specific information is given in
3 Table 1. The JM site is located in the center of the Taichung urban basin. This site is 15 m
4 above ground level (AGL) and within 150 m to a major, high-traffic crossroad of width
5 35 m. There are a total of ~ 2,610,000 registered vehicles in the Taichung City, ranking
6 3rd in the country (TW Ministry of Transportation and Communications,
7 <http://www.motc.gov.tw/>). Among these vehicles, 65% are 4-stroke and 2-stroke scooters
8 (ratio=7:3), 30% are mostly gasoline-powered passenger cars, and the remaining 5% are
9 trucks, buses and other types of vehicles. The SS site is situated 4.8 km to the western
10 coastline of Taiwan, 10 km to the east of a major industrial park, 9 km south to the largest
11 coal-fired power plant in Taiwan. This site is 13 m AGL and 150-m away from a low-
12 traffic road of width 23 m. The PL site is positioned in a relatively rural basin inside the
13 mountainous regions, where to its east is the Central Mountain Ranges of height 2-3 km.
14 An east-west direction corridor is the only major route that connects this area to the rest
15 of the Central Taiwan areas. The PL site has an elevation of 454-m above the mean sea
16 level, and is located 21-m AGL, 25-m away from the nearest major road. The JS site is
17 near the southern border of the CTAQMD, in between hills that stretch in north-south
18 direction, 44 km to the southeast of the Taichung urban area, 8-m AGL and 150-m away

1 from a major road of width 15 m. The prevailing northeasterly trade winds during winter
2 months are conducive for transport of air pollutants from the Central Taiwan region to the
3 JS site. As a result, the TW EPA had designated it as a photochemistry site.

4 **2.2 Instrumentation**

5 The number size distributions of ambient aerosols were measured by a scanning
6 mobility particle sizer (SMPS; Model 5.500, GRIMM Aerosol Technik, GmbH, Germany),
7 housed inside the monitoring sites. The sampling set up was nearly identical among all
8 the sites, described in the following. Ambient sample air was drawn through the site
9 rooftop at a flow rate of 8 liters min⁻¹ (lpm) with a 3-m glass tube. The tube of 2.5-m was
10 extended vertically outdoor and shielded inside an aluminum tube. A short stainless steel
11 forward-facing probe was inserted into the glass tube to draw a small portion of the
12 sample air into the SMPS through a 1.5-m conductive tubing at 0.3 lpm. The diffusion
13 loss to the glass and conductive sample line was estimated to be ~15% for the smallest
14 measureable particles of 11.1 nm (Baron and Willeke, 2001), though no loss corrections
15 were made to the measured data. The air conditioners inside the sites were activated only
16 when the indoor temperature exceeded 31 °C; therefore, the indoor as well as the sample
17 air temperature basically follows the ambient temperature trend. For example, the highest

1 daily ambient temperature of 31.5 °C was slightly higher than the sample air temperature
2 inside the differential mobility analyzer (DMA) of 29.5 °C on that same day. Similarly,
3 the lowest daily ambient temperature of 11.8 °C was slightly lower than the sample air
4 temperature inside the DMA of 15.4 °C on that same day. As a result, there were no
5 observed water condensation issues with the sampling tube.

6 The SMPS consists of a long Vienna-type DMA (L-DMA; Model 55-900) and a
7 butanol-based condensation particle counter (CPC; Model 5.403). The operating
8 principles of the system are described in detail elsewhere (Winklmayr et al., 1990;
9 Reischl et al., 1997; Heim et al., 2004). The inlet of the L-DMA consists of an impactor
10 with a 50% collection efficiency at 1082 nm, followed by an Am-241 neutralizer (Model
11 5.522). With its default set up, the inner electrode of the L-DMA is positively charged.
12 The detectable aerosol mobility diameters range from 11.1 to 1083.3 nm (44 size bins)
13 with a sheath and sample flow of 3 lpm and 0.3 lpm, respectively. Due to the humid
14 condition in Taiwan, the sheath air is dried to a RH of ~10% through a silica gel canister
15 and then passed through a HEPA filter before entering the L-DMA. The saturator and
16 condenser temperature of the CPC was set at 40 °C and 15 °C, respectively. The CPC has
17 a 50% counting efficiency at 4.5 nm and measures number concentration up to $2 \times 10^4 \text{ cm}^{-3}$
18 ³ with single particle counting and coincidence correction, and up to 10^7 cm^{-3} with the

1 photometric mode. The SMPS was set to down-scan 6 min and 26 s from 10,000 to 5 V
2 plus a wait-time of 34 s, producing one average particle number size distribution every 7
3 min. The SMPS has been routinely sent back to the manufacturer for calibration each year.
4 In addition, the sizing accuracy was determined before the study by sampling NaCl of
5 known mobility sizes (50, 76, 113, 168, and 241 nm) from a monodisperse aerosol
6 generator (TSI; Model 3475), simultaneously with another SMPS (TSI; Model 3936).
7 The differences between the expected and the measured mode sizes were less than 3 nm.

8 On-site air quality and meteorological data were obtained from the monitoring sites
9 and validated by the TW EPA (TW EPA, 2009). At the sites, PM₁₀ and PM_{2.5} were
10 continuously monitored by means of β -ray attenuation, SO₂ by UV fluorescence, NO_x by
11 chemiluminescence, CO by nondispersive IR, and O₃ by UV absorption. Meteorological
12 parameters used in this study include T, RH, wind speed and direction.

13 **2.3 Sampling Campaigns**

14 Two intensive sampling campaigns were carried out from October 2008 to January
15 2009 and August 2010 to October 2010. The site and sampling information are given in
16 Table 1. Based on the overall average temperatures (20.5 vs. 27.6 °C), we conveniently
17 referred to the first and second campaign as the “cold” and “warm” period, respectively

1 (Table 2). Limited by one SMPS, we continuously measured the aerosol number size
2 distributions for two weeks at each site, and rotated consecutively among the four sites.
3 Between each site, we carried out routine maintenance, including impactor cleaning,
4 system leak tests, zero-tests, and flow rate measurements. During the study periods, we
5 collected a total 26,075 aerosol number size distributions over 137 effective sampling
6 days. The overall data coverage is ~92.5%.

7 **2.4 Data Analysis**

8 The aerosol data were subjected to quality control and assurance procedures
9 described below. First, the data from the largest size bin (1083.3 nm) were discarded to
10 minimize the effect from potential multiple charged particles. Second, following the
11 screening method proposed by Yu et al. (2004), the number concentrations and the
12 relative standard deviations between particle counts of size bins below 100 nm were
13 analyzed to detect potential abnormal size distributions. Only very small numbers of
14 collected aerosol data (<0.1%) were identified as outliers that typically occurred at the
15 beginning of the instrument set up. The outliers were not included in the data analysis.
16 For convenience of discussion, the “total” submicrometer particles hereafter refer to those
17 with sizes between 10 to 1000 nm. Furthermore, the size-fractionated number

1 concentrations of nucleation (10-25 nm), Aitken (25-100 nm), accumulation (100-1000
2 nm), and ultrafine (10-100 nm) particles hereafter are denoted as N_{NUC} , N_{AIT} , N_{ACC} , N_{TOT} ,
3 and N_{UFP} , respectively. In order to match the time resolution of the air pollutant and
4 meteorological data, the aerosol data were averaged to derive hourly averages of particle
5 number size distributions. The hourly aerosol, air pollutants and meteorological data were
6 then used to compute Pearson product moment correlation coefficients (r) and to
7 determine the temporal variation. In the correlation analysis, each site's datasets for the
8 cold and warm period were combined together to cover more diverse pollutant levels and
9 meteorological conditions as well as increasing the sample size.

10 The mean particle number size distributions for each site and study period were
11 fitted with three lognormal distributions to identify the underlying sub-modes. The
12 selection of tri-modal fits is because bimodal fits, in a number of cases, failed to account
13 for the accumulation particles. The fitting process was done using the following
14 distribution function (df):

$$df = A \exp \left[- \left(\frac{(\ln D_p - \ln D_m)}{W} \right)^2 \right] \quad (1)$$

15 where A sets the amplitude (peak height), D_m sets the particle mode diameter, and W sets
16 the peak width. The Levenberg-Marquardt algorithm, a nonlinear least-squares fitting, is

1 used to search for the minimum value of chi-square (χ^2) (Igor Pro v.6.2.2.2, WaveMetrics,
2 Inc.). The square root of the number concentration for each size bin (i.e., the
3 measurement uncertainty of particle counter) was used in the calculation of χ^2 to obtain a
4 more accurate fit as follows:

$$\chi^2 = \sum \left(\frac{y - y_i}{w_i} \right)^2 \quad (2)$$

5 where y is the fitted value, y_i is the measured value, and w_i is the measurement
6 uncertainty for size bin i . The tri-modal fitting terminates after 40 passes in searching for
7 the best fit, or will quit if 9 passes in a row do not result in a decrease in χ^2 .

8 **3. Results and Discussion**

9 **3.1 Meteorological Conditions and Air Quality**

10 A detailed description of the typical synoptic weather patterns for each season in
11 Taiwan has been given elsewhere (e.g., Cheng, 2001). The site-specific meteorological
12 conditions during the study are summarized in Table 2. The temporal evolution of the
13 wind speed at each site during the two study periods is presented in Figure S1. With
14 respect to average values, the major differences between the cold and warm period
15 include the ambient temperature (20.6 vs. 27.6 °C), prevailing wind direction (NNW vs.
16 W) and accumulated precipitation (57.4 vs. 283.6 mm). As Taiwan is situated at the

1 border of the subtropical region, it is important to note that the meteorological differences
2 between seasons are relatively small compared to mid- or high-latitude regions. For
3 example, the minimum and maximum monthly average temperature over the past three
4 decades were 16.6 °C (January) and 28.6 °C (July), respectively, i.e., a range of only 12
5 °C (Taiwan Central Weather Bureau, <http://www.cwb.gov.tw/>). Based on these historical
6 data, it shows that the present cold period was warmer ($\Delta T=4$ °C), whereas the warm
7 period is comparable ($\Delta T=1$ °C) to typical conditions.

8 The site-specific hourly averages and diurnal patterns of air pollutant concentrations
9 are given in Table 3 and Figure S2, respectively. Overall, the major air quality differences
10 between the cold and warm period include the significantly higher PM and NO_x, and
11 slightly higher SO₂ concentration during the cold period. Furthermore, despite the large
12 seasonal and site-to-site differences in PM concentrations, the average percentage of
13 PM_{2.5} over PM₁₀ of 65% remained nearly constant.

14 **3.2 Submicrometer Particle Number Concentrations**

15 The summary statistics of the hourly N_{NUC} , N_{AIT} , N_{ACC} , and N_{TOT} and the fraction of
16 $N_{\text{UFP}}/N_{\text{TOT}}$ during the two study periods are summarized in Table 4. The highest mean
17 N_{TOT} of $3.1 \times 10^4 \text{ cm}^{-3}$ and $3.4 \times 10^4 \text{ cm}^{-3}$ were measured at the urban site for the cold and

1 warm period, respectively. These mean N_{TOT} were comparable to that measured at a
2 number of other urban sites, in street canyon and at roadside, but lower than that
3 measured on road and inside tunnels (Stanier et al., 2004b; Harrison and Jones, 2005;
4 Aalto et al., 2005; Moore et al., 2009; Lonati et al., 2011; Bae et al., 2010; Asmi et al.,
5 2011; Wang et al., 2010; Chen et al., 2010a; Cheng et al., 2010a; Cheng et al., 2010b;
6 Wang et al., 2011b; Sabaliauskas et al., 2011). The mean N_{TOT} at the coastal, mountain,
7 and downwind site were similar, in the range of 1.7×10^4 - 2.2×10^4 cm^{-3} . These averages are
8 29-49% lower than that measured at the urban site, attributable to the substantially higher
9 N_{UFP} (i.e., $N_{NUC} + N_{AIT}$) at the urban site. This observation is consistent with earlier studies
10 that showed higher number concentrations at sites impacted by heavy traffic, whereas
11 lower concentrations were distributed more or less homogeneously at background or
12 receptor sites (Harrison and Jones, 2005; Hudda et al., 2010; Bae et al., 2010). Another
13 explanation to the higher N_{UFP} is the frequent NPF events observed at the urban site
14 during the warm period, discussed later in Section 3.3. On the other hand, excluding the
15 storm-impacted downwind site, the lowest mean N_{TOT} of 1.7×10^4 - 1.8×10^4 cm^{-3} were
16 measured at the mountain site during both the cold and warm periods. The average $\text{PM}_{2.5}$
17 levels (up to $47.7 \mu\text{g m}^{-3}$; Table 3) were typically highest at the mountain site.
18 Additionally, its $\text{PM}_{2.5}$ and N_{UFP} were negatively correlated ($r = -0.25$; see Section 3.5).

1 This suggests a weaker UFP source strength and/or the elevated pre-existing particles are
2 not conducive for the survival of the UFP at the mountain site. Modeling studies have
3 shown that nucleation particles are quickly removed by coagulation with larger
4 background particles or by condensational growth, with time scales of as small as 0.01-
5 0.1 hr (Pandis et al., 1995; Kerminen et al., 2004).

6 On a daily basis, the mean N_{TOT} on weekdays and weekends were $2.3 \times 10^4 \text{ cm}^{-3}$ and
7 $2.2 \times 10^4 \text{ cm}^{-3}$, respectively, with a slightly higher variability during the weekend (Table
8 S1). As shown, the weekday and weekend N_{TOT} differences during the cold and warm
9 period were negligible. Asmi et al. (2011) also have shown no statistically significant
10 variation of concentrations between weekdays and weekends in a number of European
11 sites. On a seasonal basis, the overall mean N_{TOT} at each site during the warm period was
12 generally higher than that during the cold period. The site-specific seasonal differences
13 were $< 22 \%$. This shows that the weekday-to-weekend as well as the cold-to-warm
14 seasonal variability is smaller than the site-to-site spatial variability of N_{TOT} . Unlike the
15 present study, many others have shown a distinct seasonal pattern and weekday-weekday
16 differences in the particle number concentrations, with higher levels observed in colder
17 months and during weekdays (Singh et al., 2006; Aalto et al., 2005; Stanier et al., 2004b;
18 Harrison and Jones, 2005; Moore et al., 2009; Birmili et al., 2010; Bae et al., 2010;

1 Lonati et al., 2011; Wang et al., 2011a; Sabaliauskas et al., 2011). Paatero et al. (2005)
2 and Sabaliauskas et al. (2011) have shown a negative correlation between particle number
3 and T, where the later covers a range from -4.6 to 28 °C. Marawska et al. (2008)
4 suggested that the seasonal variability is more distinct in mid- and high-latitude regions in
5 the Northern Hemisphere, where the meteorological differences between seasons are
6 significant. For example, no seasonal pattern was identified for the particle number
7 concentration in subtropical Brisbane, Australia (Mejia et al., 2007). Similarly, strong
8 seasonal variability is considered unlikely due to the small temperature difference (~7 °C)
9 between the present two study periods.

10 Within the submicrometer size range, the percentages of $N_{\text{UFP}}/N_{\text{TOT}}$ were in the
11 range of 69% to 87%, between which higher values were observed at the urban and
12 coastal site (Table 4). These observations are comparable to that in Milan, Italy (Lonati et
13 al., 2011), but lower than that of 90% in Los Angeles, CA (Moore et al., 2009). The site-
14 to-site as well as the cold-to-warm percent differences of $N_{\text{UFP}}/N_{\text{TOT}}$ were <17%. Given
15 that the overall mean N_{TOT} of $2.2 \times 10^4 \text{ cm}^{-3}$, the N_{UFP} differences were $<0.4 \times 10^4 \text{ cm}^{-3}$. The
16 overall averages and variability of $N_{\text{UFP}}/N_{\text{TOT}}$ for the cold and warm period were nearly
17 identical. Overall, the N_{NUC} , N_{AIT} , and N_{ACC} were generally similar among the four sites
18 and between the cold and warm period. An exception is the considerably higher N_{NUC} and

1 N_{AIT} in the range of 1.1×10^4 - 1.6×10^4 cm^{-3} at the urban site than at the other three sites.
2 This indicates that the urban area is an important source of freshly-formed UFP. On the
3 other hand, the coastal, mountain and downwind site have comparable, relatively weaker
4 UFP source strength, and that their atmospheric conditions (e.g., high PM levels at the
5 mountain and downwind site or strong wind speed at the coastal site) may be unfavorable
6 for the survival of local UFP.

7 **3.3 Submicrometer Particle Number Size Distributions**

8 The temporal evolution of the particle number size distributions measured at the four
9 sites during the cold and warm period is presented in Figure 2. A key feature common to
10 all the sites is that the particle number size distributions show strong diurnal variability
11 during both study periods. The variability is characterized by the active production of
12 UFP during the daytime hours and the disappearance of nucleation particles during the
13 midnight to early morning hours. This indicates that the human activities and
14 photochemistry are the two dominant driving forces behind the emission or production of
15 UFP. Such UFP production is particularly active at the urban site. In specific, we
16 identified frequent NPF events (eight out of 12 days) during the warm period of $\text{PM}_{2.5}$
17 averaging $21.2 \mu\text{g m}^{-3}$. Notably, such $\text{PM}_{2.5}$ levels are considered “low” in Taiwan (see

1 Table 3), but not so at urban sites in other countries (e.g., Bae et al., 2010; Wang et al.,
2 2011b). The NPF events are characterized by a burst of the smallest measurable particles
3 of 11.1 nm, followed by continuous particle growth over a period of up to 5 hours.
4 Furthermore, the onset of the urban NPF events was typically before 09 local time (LT),
5 not during midday hours. This suggests the NPF was closely connected to the morning
6 traffic emissions and possibly the breakup of nocturnal inversion layer. Wang et al.
7 (2011a) also reported frequent morning nucleation events in Rochester, NY. However,
8 unlike the present study, those events were not accompanied by particle growth. There
9 were also several NPF events at the other sites, though the frequency and intensity were
10 much lower. At the coastal site, the onset of UFP production and the number size
11 distribution characteristics were more variable than the other sites. At the mountain and
12 downwind site, the onset of UFP production was relatively consistent during the
13 afternoon or midday hours. However, the UFP production was substantially more intense
14 at the downwind site. Details of the NPF and growth events at the four sites will follow in
15 subsequent publication.

16 The statistical presentations (16th, 50th, 84th percentiles and mean) of the site-
17 specific particle number size distributions are shown in Figures 3 and 4, respectively.
18 Also shown in the figures are the tri-modal lognormal fits to the mean particle number

1 size distributions. The site-specific size distributions showed considerable temporal
2 variability. In specific, the mean concentrations were typically higher than the 50th
3 percentile concentrations (i.e., skew to the right) for particles in the UFP size range. All
4 the size distributions exhibited two dominant modes, namely the nucleation and Aitken
5 mode, and a minor accumulation mode. Such tri-modality is in agreement with that
6 observed in Milan, Italy (Lonati et al., 2011), but in contrast to the uni-modal median size
7 distributions measured at a number of European sites (Asmi et al., 2011). In this study,
8 the location and intensity (i.e., size and concentration) of the modes show considerable
9 site-to-site and cold-to-warm variability. Nevertheless, three common features are
10 summarized as follows. First, a nucleation mode is always present in the site-specific
11 mean size distributions for the cold and warm period. This indicates the omnipresence of
12 freshly-formed particles. Second, the Aitken and accumulation mode sizes are always
13 smaller during the cold period. This feature may be related to the relatively lower
14 temperature (Table 2), which favors the formation of fresh and smaller particles.
15 Sabaliauskas et al. (2011) observed the geometric mean diameter was smallest during the
16 winter due to increased concentrations of <50 nm particles. Third, the concentrations of
17 the accumulation mode particles are always considerably higher during the cold period.
18 This feature is consistent with the significantly higher PM levels during the cold period

1 (Table 3). Wang et al. (2011b), however, reported greater accumulation-mode
2 contributions to the total particle numbers in the summer due to regional transport.
3 Overall, the results show that there is notable variability in the particle number size
4 distributions among sites and between study periods, despite having comparable N_{TOT} .
5 This implies that the total particle number concentration alone may not be a
6 representative exposure or impact estimate, both of which depend on particle size.

7 **3.4 Diurnal Variations of Submicrometer Particles**

8 The diurnal variability of the hourly N_{TOT} and air pollutants concentrations at each
9 site is shown in Figures 5 and S2, respectively. During the cold period, the N_{TOT} at the
10 urban, downwind and mountain site showed distinct and consistent diurnal patterns with
11 peak concentrations between 08-09 and 17-19 LT. The peak concentrations ranged from
12 $1.6 \times 10^4 \text{ cm}^{-3}$ to $4.4 \times 10^4 \text{ cm}^{-3}$, between which higher values were measured at the urban
13 site and lower ones at the mountain site. Many other studies have shown such a daily
14 two-peak pattern, particularly in the urban area and on workdays (Wehner and
15 Wiedensolher, 2003; Jeong et al., 2004; Hussein et al., 2004; Stanier et al., 2004b;
16 Harrison and Jones, 2005; Birmili et al., 2010). The temporal patterns of CO and NO_x , to
17 a lesser extent for SO_2 , correlated well with that of N_{TOT} , especially during the morning

1 traffic rush hours. This indicates that those particles were predominantly from traffic
2 emissions. Unlike others, the temporal patterns of the size-fractionated particles at the
3 coastal site were quite different from the three other sites. For example, the morning peak
4 concentration at the coastal site was less pronounced and there was no obvious evening
5 peak during the cold period. This is likely due to the considerably higher average wind
6 speeds of 4.3 m s^{-1} near the coast, conducive for atmospheric dispersion. A closer
7 examination of the size-fractionated particle number concentrations reveals that their
8 temporal behaviors are dependent on particle size during the cold period (Figure 6). As
9 shown, the production of nucleation particles was more active in the morning, at noon
10 and in the early evening. The Aitken mode particles were driving the N_{TOT} variability,
11 evident by their nearly identical temporal patterns. The accumulation particles generally
12 followed the temporal patterns of Aitken mode particles, though there is an obvious
13 impact of aged particles at the downwind site. This is likely due to the winter prevailing
14 northerly winds that favors the transport of pollutants from the upwind urban area (Table
15 2 and Figure 1).

16 During the warm period, the temporal pattern of N_{TOT} at the four sites was more
17 variable than that during the cold period (Figure 5). The diurnal variability of the hourly
18 N_{NUC} , N_{AIT} and N_{ACC} at each site during the warm period is given in Figure 7. The

1 mountain site is the only site where the traffic-related two-peak concentration profile of
2 N_{TOT} remained visibly distinct. This is likely because of its unique topography, where the
3 surrounding Central Mountain Ranges are favorable for trapping the air mass that enters
4 through the east-west direction corridor. In the contrary, there were no obvious traffic-
5 related morning N_{TOT} peak concentrations at the urban, coastal and mountain site. In
6 particular, the urban site showed a distinct peak concentration of $6.9 \times 10^4 \text{ cm}^{-3}$ at 10 LT.
7 This peak was not only considerably more pronounced but also occurred an hour later
8 than the typical traffic-related morning peak. It was due to the frequent, traffic-related
9 NPF and growth events in the morning, evident by the substantially elevated N_{NUC} and
10 N_{AIT} shown in Figure 7. In comparison, three other intense NPF and growth events were
11 observed only at the downwind site during the warm period. At the coastal and downwind
12 site, the diurnal pattern showed relatively elevated concentrations between 11-15 LT.
13 Compared to the urban site, these daytime peak concentrations were considerably less
14 intense, and they overlapped more during periods of strong solar radiation. This is
15 suggestive of the link between photochemistry and NPF. The more intense and frequent
16 NPF at the urban site is likely a result of its higher traffic density, hence higher
17 condensable vapor concentrations (e.g., H_2SO_4 and organics). Overall, the nucleation and
18 Aitken mode particles showed distinctly different temporal patterns between the cold and

1 warm period. The accumulation particles during the warm period showed less consistent
2 patterns, except the obvious impact of aged particles at the mountain site due to the
3 summer prevailing westerly winds (Table 2). Winds from that direction likely inhibited
4 the dispersion of air pollutants in the mountainous region by blocking the outward
5 easterly airflow (Figure 1).

6 Regardless of the study periods and sites, the average N_{TOT} typically reached a
7 minimum of $\sim 1.0 \times 10^4 \text{ cm}^{-3}$ in the early morning hours during 03-06 LT, of which could
8 be regarded as the ambient background particle number concentration in the study area.
9 Such background level is consistent with that measured at the roadside in Taiwan (Chen
10 et al., 2010a).

11 **3.5 Size-Fractionated Particle Numbers, Copollutants and Meteorology**

12 The correlation coefficients (r) between the hourly size-fractionated particle
13 numbers, selected copollutants and meteorological parameters at each site are given in
14 Table 5. As the $PM_{2.5}$ correlated very well with the PM_{10} ($r = 0.85-0.94$), hence the latter
15 is not shown in Table 5. The N_{NUC} at the urban site showed negligible to poor correlations
16 with all the copollutants ($r < 0.11$), weak negative correlation with RH ($r = -0.33$) and
17 positive correlation with wind speed (WS) ($r = 0.17$). This indicates that the formation of

1 nucleation particles at the urban site is not solely related to primary emissions (CO, NO_x
2 and SO₂), and that it is favored under dry and increased wind speed (i.e., daytime hours
3 and enhanced mixing; Figure S1). In downtown Toronto, Canada, Sabaliauskas et al.
4 (2011) however showed the nucleation particles were not correlated with solar radiation
5 and SO₂ but likely better predicted by NO_x. At the coastal and downwind site, the N_{NUC}
6 were positively correlated with SO₂ (r = 0.29 and 0.33, respectively), negatively
7 correlated with RH (r = -0.48 and -0.37, respectively) and positively correlated with WS
8 (r = 0.18 and 0.19, respectively). However, the N_{NUC} at the downwind site was also
9 positively correlated with O₃, indicating that the secondary photochemical processes are
10 possibly an important influential factor. A number of studies have reported NPF events
11 are strongly tied to the photochemistry under the presence of elevated SO₂ (e.g.,
12 Verheggen and Mozurkewich, 2002; Stanier et al., 2004a; Young and Keeler, 2007; Wu et
13 al., 2007). Unlike others, the N_{NUC} at the mountain site was moderately correlated with
14 NO_x and CO (r = 0.49 and 0.40, respectively), suggesting a notable connection with
15 primary emissions from local traffic.

16 The N_{AIT} showed moderate correlations with NO_x and CO (r = 0.46-0.68) at all but
17 the downwind site, where it was more correlated with SO₂ (r = 0.39). This clearly
18 indicates an important connection between the Aitken mode particles and primary

1 emissions. Earlier studies have shown that the traffic-related particle number emissions
2 are dominated by the UFP fraction, where the mode diameter is highly dependent on the
3 vehicle type and distance to the road (e.g., Zhu et al. 2002; Ogulei et al., 2007; Lonati et
4 al., 2011). At the urban and mountain site, the N_{AIT} were more or less independent of the
5 T, RH and WS ($|r| < 0.15$). At the coastal site, elevated N_{AIT} were related to stagnant (WS;
6 $r = -0.40$), warm (T; $r = 0.28$) and low O_3 ($r = -0.28$) conditions. At the downwind site,
7 elevated N_{AIT} were related to conditions of elevated SO_2 ($r = 0.39$), O_3 ($r = 0.12$), WS ($r =$
8 0.22) and low RH ($r = -0.27$). This indicates the positive influence of regional transport
9 and photochemistry under dry conditions on the production of Aitken mode particles at
10 the downwind site.

11 The N_{ACC} at all sites showed moderate to strong positive correlations with all the
12 copollutants SO_2 , CO, NO_x and $PM_{2.5}$ ($r = 0.44-0.86$), except that it was not at all
13 correlated with the SO_2 at the mountain site ($r = 0.09$). This suggests that these pollutants
14 share a common source, most likely traffic emissions. The T, RH and WS generally
15 showed weak correlations with the N_{ACC} . Some exceptions include elevated N_{ACC} at the
16 downwind site were moderately related to low temperature, which are indicative of
17 particle buildup in the early evening or nighttime hours (Figure 6). In addition, elevated
18 N_{ACC} at the urban and coastal site were associated with stagnant conditions, suggesting

1 the effects of poor dispersion as well as their local-origin. The strong correlation between
2 the N_{ACC} and $PM_{2.5}$ ($r = 0.59-0.86$) is in part because a significant fraction of fine particle
3 mass is composed of accumulation particles. For example, Li and Lin (2002) and Lin et al.
4 (2005) have shown that the average $PM_1/PM_{2.5}$ ratio were above 0.76 at urban and traffic
5 sites in Taiwan. Wang et al. (2011a) also observed a strong and stable association between
6 the N_{ACC} and $PM_{2.5}$. In contrast, the N_{NUC} and N_{AIT} showed weak correlations with $PM_{2.5}$
7 ($|r| < 0.25$). This suggests that the UFP and fine particles have different sources or
8 formation processes, particularly the nucleation particles. For example, the correlation
9 between N_{NUC} and $PM_{2.5}$ ($r = -0.25$) was negative at the mountain site and negligible at
10 the three other sites ($|r| < 0.09$). The former in particular shows that elevated $PM_{2.5}$ could
11 be unfavorable for the production or existence of the nucleation particles. Nevertheless,
12 the Aitken mode and accumulation particles at all of the study sites are predominantly
13 related to primary emissions from traffic. The weak correlations but common traffic-
14 origin between the UFP and accumulation particles underscore that each of the two
15 particle fractions has another source not shared by the other, and/or they have different
16 loss processes in the atmosphere. For example, the secondary photochemical pathway
17 appears to be more relevant to the production of UFP than to that of accumulation
18 particles in the present study area. Furthermore, the much longer atmospheric lifetime of

1 the accumulation particles allows them to survive under polluted environments, in which
2 the UFP are quickly scavenged by pre-existing particles, as illustrated at the mountain site.

3 **4. Conclusion**

4 This study presents first measurements of the submicrometer (10-1000 nm) particle
5 number size distributions at four distinct types of environment (urban, coastal, mountain
6 and downwind) within the Central Taiwan Air Quality Management District during a cold
7 (20.5 °C) and a warm (27.6 °C) period. The urban site showed the highest average
8 submicrometer particle number concentration of up to $3.4 \times 10^4 \text{ cm}^{-3}$, which is 29 to 49%
9 higher than the coastal, mountain and downwind site. The latter three sites showed
10 comparable number concentration, ranging from 1.7×10^4 to $2.2 \times 10^4 \text{ cm}^{-3}$. The average
11 submicrometer particle number concentration at each site showed small variability (<
12 22%) between the cold and warm period, despite notable differences in the
13 meteorological conditions and air quality. During the cold period, the hourly
14 submicrometer particle number concentrations showed distinct diurnal patterns at all but
15 the coastal site with peak concentrations coincided with rush-hour traffic emissions (CO
16 and NO_x). During the warm period, the diurnal patterns across the four sites were more
17 variable and the influence of traffic emissions was relatively weaker. Furthermore, the
18 photochemistry and SO₂ possibly from combustion sources or regional transport were

1 associated the daytime peak concentrations. Despite having similar submicrometer
2 particle number concentrations, the modal characteristics of the measured particle number
3 size distributions exhibited notable spatial and temporal differences. The ultrafine (10-
4 100 nm) particles and the accumulation (100-1000 nm) particles correlated poorly with
5 each other, even though both were closely connected to traffic emissions. In addition, the
6 accumulation particles were highly correlated with $PM_{2.5}$ and PM_{10} . Such disconnection
7 was likely due to the influences of sources other than road traffic and/or the differences in
8 their ambient loss processes in different environments. Overall, the results indicate
9 modest spatial heterogeneity and strong diurnal variability of the submicrometer particle
10 number concentrations. In addition, there are variable sources or meteorology-dependent
11 formation processes of the ultrafine particles within the study area. In view of the above
12 variability, measurements of the “total” particle number concentrations without
13 discriminating sizes or using PM mass concentrations to estimate particle numbers at a
14 single site would be inadequate for exposure and impact assessment in a region of diverse
15 environments. Instead, at a minimum, the number concentrations of ultrafine and
16 accumulation particles should be measured separately at different types of environment in
17 order to more accurately determine their spatiotemporal variability and hence impact.

1 **Acknowledgments**

2 The authors wish to express their appreciation to the Taiwan EPA for making the air
3 quality monitoring sites available for the present study. The financial support from the
4 Taiwan National Science Council (NSC-97-2218-E-039-002-MY3) and the China
5 Medical University (CMU-96-109) are gratefully acknowledged.

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- 4 of ultrafine particles near major highways in Los Angeles. *Aerosol Sci Technol*
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2 **TABLE AND FIGURE CAPTIONS**

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4 **Table 1.** Site and sampling information during the entire study.

5 **Table 2.** Hourly averages of meteorological conditions at each site during the two study

6 periods.

7 **Table 3.** Hourly averages of air pollutant concentrations at each site during the two study

8 periods.

9 **Table 4.** Summary statistics of the hourly size-fractionated number concentrations (10^4

10 cm^{-3}) at each site during the two study periods.

11 **Table 5.** Correlation coefficients between size-fractionated particle number, air pollutant

12 concentrations and meteorological parameters at each site.

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2 **Fig. 1.** The urban, coastal, mountain and downwind site under study in the Central
3 Taiwan Air Quality Zone.

4 **Fig. 2.** The temporal evolution of the measured particle number size distributions at the (a)
5 urban, (b) coastal, (c) mountain and (d) downwind site during the cold and warm
6 period.

7 **Fig. 3.** Statistical presentations and ti-modal lognormal fits of the particle number size
8 distributions at the (a) urban, (b) coastal, (c) mountain and (d) downwind site
9 during the cold period.

10 **Fig. 4.** Statistical presentations and tri-modal lognormal fits of the particle number size
11 distributions at the (a) urban, (b) coastal, (c) mountain and (d) downwind site
12 during the warm period.

13 **Fig. 5.** The diurnal variability of the hourly submicrometer particle number
14 concentrations (N_{TOT}) at each site during the (a) cold and (b) warm period.

15 **Fig. 6.** The diurnal variability of the hourly (a) N_{NUC} , (b) N_{AIT} and (c) N_{ACC} at each site
16 during the cold period.

17 **Fig. 7.** The diurnal variability of the hourly (a) N_{NUC} , (b) N_{AIT} and (c) N_{ACC} at each site
18 during the warm period.

Table 1. Site and sampling information during the entire study.

Site code	Type	Elev. (m)	Sampling height (m)	Latitude (North)	Longitude (East)	Sampling duration		Aerosol samples ^a
						Cold period	Warm period	
JM	Urban	95	16	24°09'31"	120°39'34"	10/3-10/20/2008, 12/13-12/25/2008	8/9-8/20/2010	8293
SS	Coastal	6	19	24°07'54"	120°28'8"	10/29-11/10/2008, 12/29/08-1/9/2009	9/15-9/24/2010	6763
PL	Mountain	454	17	23°58'8"	120°58'4"	11/17-11/28/2008, 1/9-1/22/2009	10/11- 10/22/2010	5836
JS	Downwind	8	12	23°45'23"	120°40'38"	11/28-12/12/2008	9/2-9/13/2010	5183

a. The time resolution of SMPS measurements was 7-min.

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Table 2. Hourly averages of meteorological conditions at each site during the two study periods.

	n ^a	T (°C)	RH (%)	WS (m s ⁻¹)	Prevailing WD	Rain ^b (mm)
Cold period (Oct.-Jan.)						
Urban	657	22.8	70	0.6	NNE	22.2
Coastal	562	21.5	78	4.3	NNW	33.2
Mountain	436	18.2	70	1.0	WSW	1.6
Downwind	356	17.4	66	0.8	N	0.4
Overall	2011	20.5	71	1.8	NNW	57.4
Warm period (Aug.-Oct.)						
Urban	288	30.3	72	1.1	S	55.6
Coastal	240	27.2	81	2.4	NE	22.2
Mountain	288	25.9	75	1.1	W, WNW	13.0
Downwind ^c	288	26.9	83	1.1	SW, WSW	192.8
Overall	1104	27.6	78	1.4	W	283.6

a. Sample size.

b. Accumulated rain.

c. The downwind site was impacted by a severe tropical storm on September 9 and 10, 2010.

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Table 3. Hourly averages of air pollutant concentrations at each site during the two study periods.

Cold period (20.5 °C)		n	SO ₂	CO	O ₃	NO _x	PM _{2.5}	PM ₁₀	PM _{2.5} /PM ₁₀
Urban		652	3.9	0.7	24.6	34.1	37.9	54.6	69
Coastal		559	4.4	0.3	29.8	15.6	24.6	44.4	55
Mountain		417	3.6	0.5	25.7	20.9	47.7	66.7	72
Downwind		351	2.7	0.5	27.7	22.1	42.8	70.3	61
Overall		1979	3.8	0.5	26.9	23.9	37.2	57.1	65
Warm period (27.6 °C)									
Urban		260	3.2	0.5	24.8	22.4	21.2	31.6	67
Coastal		240	4.2	0.3	27.8	15.4	25.5	44.5	57
Mountain		245	2.2	0.4	34.4	15.0	36.9	49.3	75
Downwind		288	2.2	0.2	25.5	9.2	13.6	23.9	57
Overall		1033	2.9	0.4	28.1	15.4	24.2	37.0	65

Note: CO in ppm, PM in $\mu\text{g m}^{-3}$, and all others in ppb; PM_{2.5}/PM₁₀ in %.

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Table 4. Summary statistics of the hourly size-fractionated number concentrations (10^4 cm^{-3}) at each site during the two study periods.

Cold period (20.5 °C)	n	N _{TOT}		N _{NUC}		N _{AIT}		N _{ACC}		N _{UFP} /N _{TOT} ^a	
		Mean	SD	Mean	SD	Mean	SD	Mean	SD	Mean	SD
Urban	641	3.1	1.3	1.1	0.7	1.5	0.7	0.5	0.3	0.81	0.08
Coastal	548	1.7	0.9	0.6	0.5	0.8	0.4	0.3	0.2	0.81	0.10
Mountain	391	1.7	0.8	0.4	0.3	0.8	0.5	0.5	0.2	0.70	0.11
Downwind	331	2.2	1.2	0.5	0.4	1.0	0.6	0.6	0.4	0.69	0.11
Overall	1911	2.2	1.3	0.7	0.6	1.1	0.7	0.5	0.3	0.77	0.11
Warm period (27.6 °C)											
Urban	268	3.4	1.8	1.4	1.2	1.6	0.8	0.4	0.2	0.87	0.08
Coastal	215	2.0	1.3	0.6	0.8	1.0	0.6	0.4	0.3	0.76	0.13
Mountain	260	1.8	0.7	0.4	0.3	0.9	0.4	0.5	0.2	0.70	0.10
Downwind	264	1.7	1.0	0.5	0.4	0.9	0.7	0.3	0.1	0.80	0.10
Overall	1007	2.2	1.4	0.7	0.9	1.1	0.7	0.4	0.2	0.78	0.12

a. The fraction of ultrafine particle (N_{NUC}+N_{AIT}) to total particle (N_{TOT}) number concentration.

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Table 5. Correlation coefficients between size-fractionated particle number, air pollutant concentrations and meteorological parameters at each site.

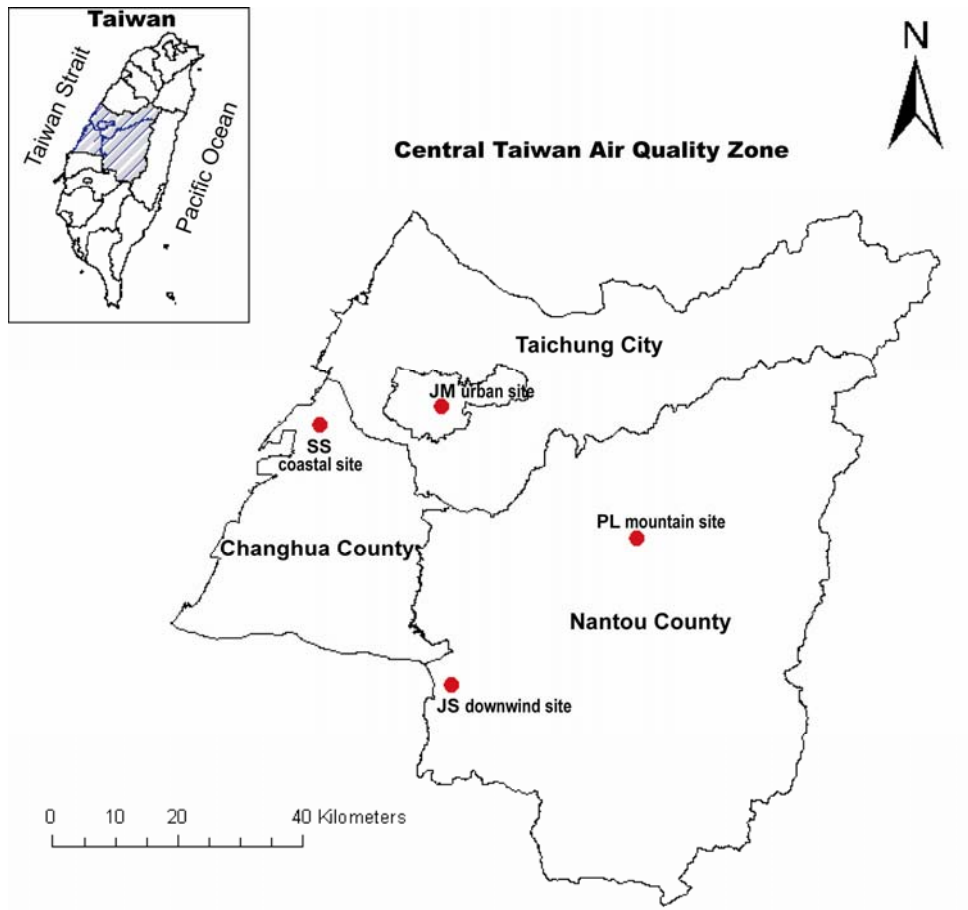
Urban	SO ₂	CO	NO _x	O ₃	PM _{2.5}	T	RH	WS
N _{NUC}	0.11	0.08	0.06	0.10	-0.08	0.16	-0.33	0.17
N _{AIT}	0.26	0.62	0.54	-0.05	0.20	0.14	-0.11	-0.08
N _{ACC}	0.45	0.77	0.73	-0.04	0.72	0.15	0.03	-0.31
Coastal								
N _{NUC}	0.29	0.07	0.10	-0.01	-0.05	-0.15	-0.48	0.18
N _{AIT}	0.44	0.46	0.58	-0.28	0.19	0.29	0.01	-0.40
N _{ACC}	0.53	0.62	0.57	-0.26	0.78	0.12	0.03	-0.39
Mountain								
N _{NUC}	0.09	0.40	0.49	-0.01	-0.25	0.15	-0.06	-0.02
N _{AIT}	0.13	0.57	0.68	0.04	-0.05	0.15	0.05	-0.09
N _{ACC}	0.01	0.62	0.49	0.11	0.59	0.14	0.13	-0.14
Downwind								
N _{NUC}	0.33	0.21	0.17	0.28	0.09	0.12	-0.37	0.19
N _{AIT}	0.39	0.29	0.26	0.32	0.25	0.16	-0.27	0.22
N _{ACC}	0.44	0.66	0.62	0.12	0.86	-0.42	-0.27	-0.12

Bold numbers are the highest correlation coefficients between each particle size fraction and copollutants.

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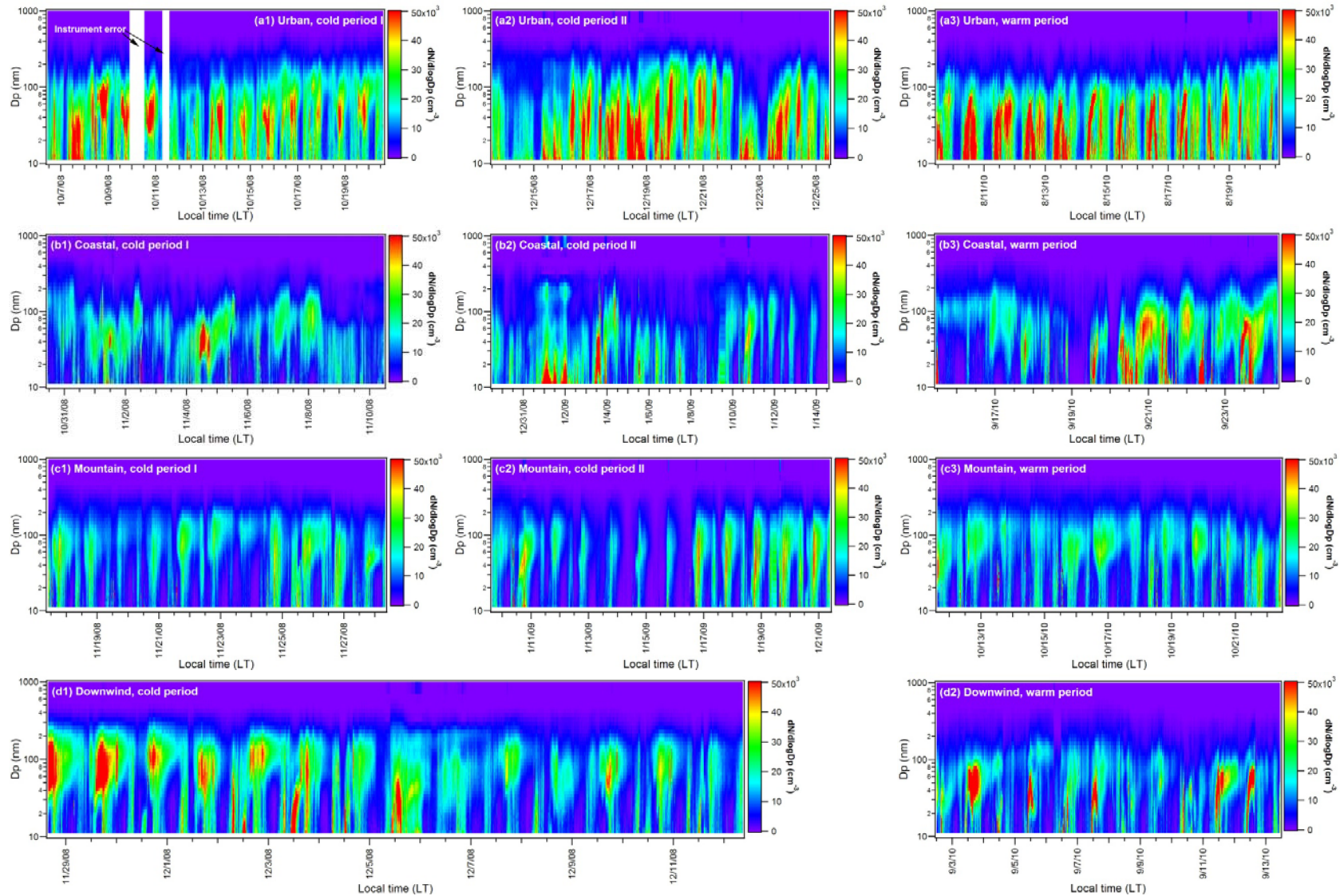
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4 **Fig. 1.** The urban, coastal, mountain and downwind site under study in the Central

5 Taiwan Air Quality Zone.



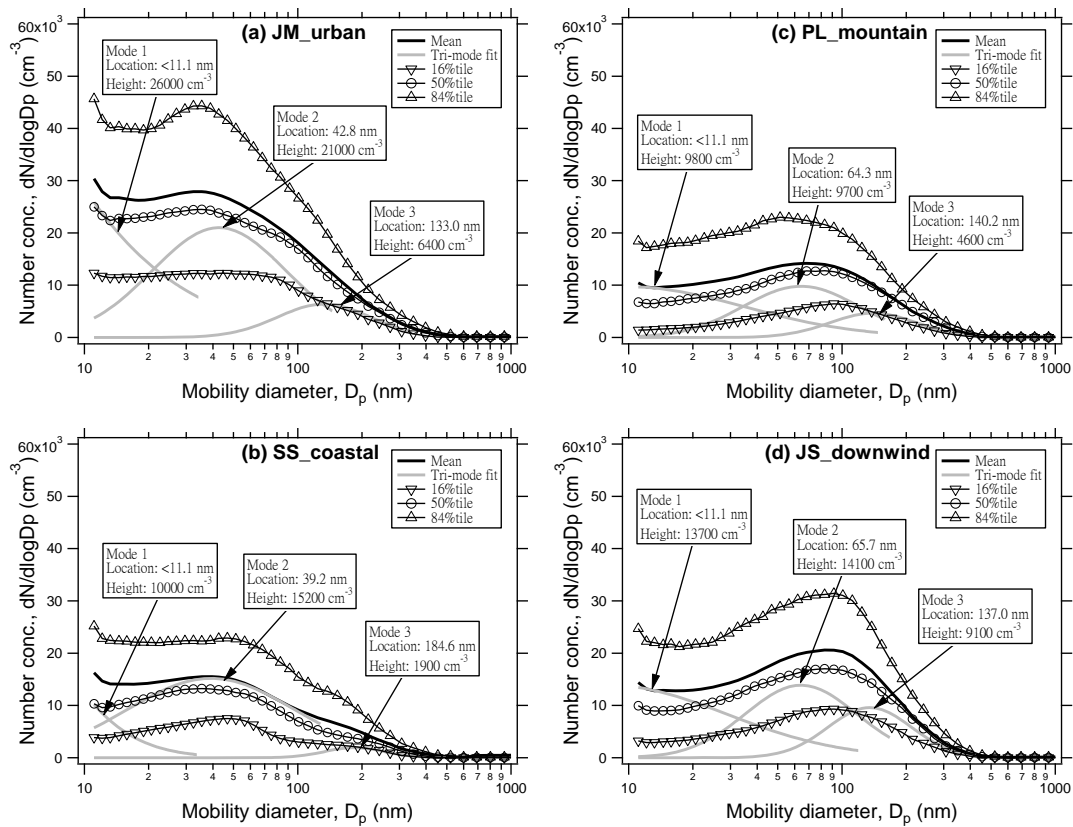
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5 **Fig. 2.** The temporal evolution of the measured particle number size distributions at the (a) urban, (b) coastal, (c) mountain and (d)
 6 downwind site during the cold and warm period.



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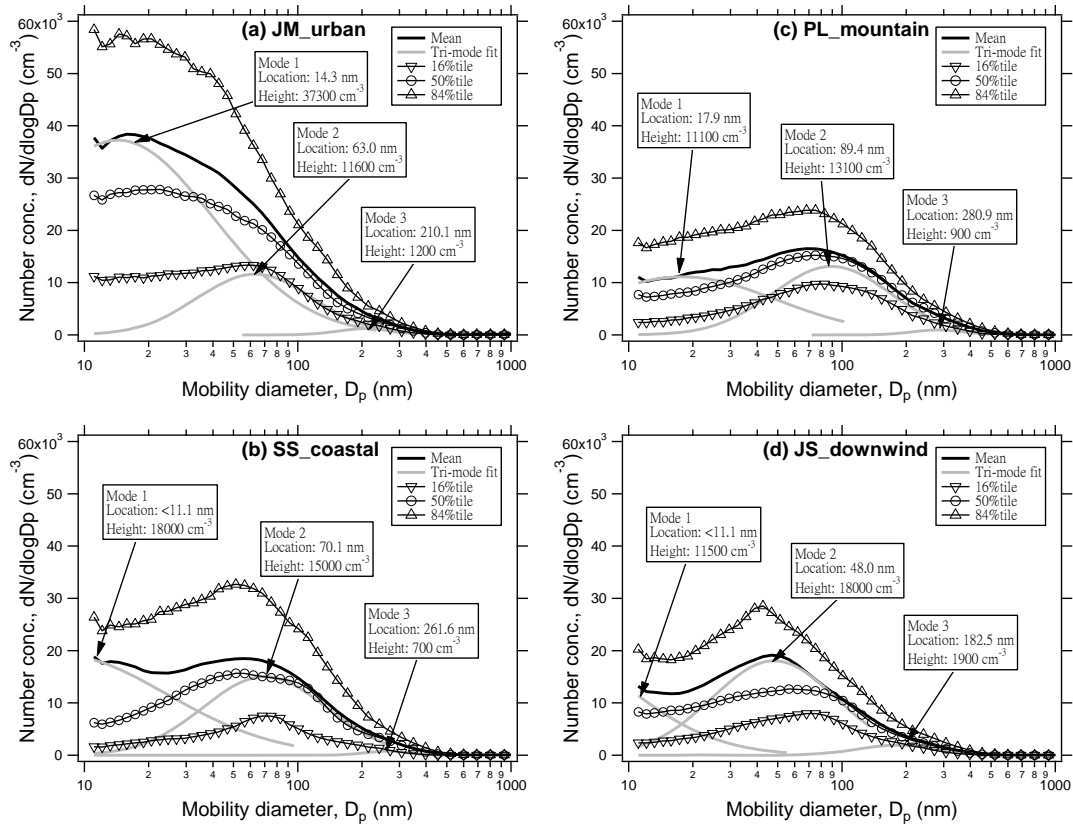
3 **Fig. 3.** Statistical presentations and tri-modal lognormal fits of the particle number size

4 distributions at the (a) urban, (b) coastal, (c) mountain and (d) downwind site during the

5 cold period.

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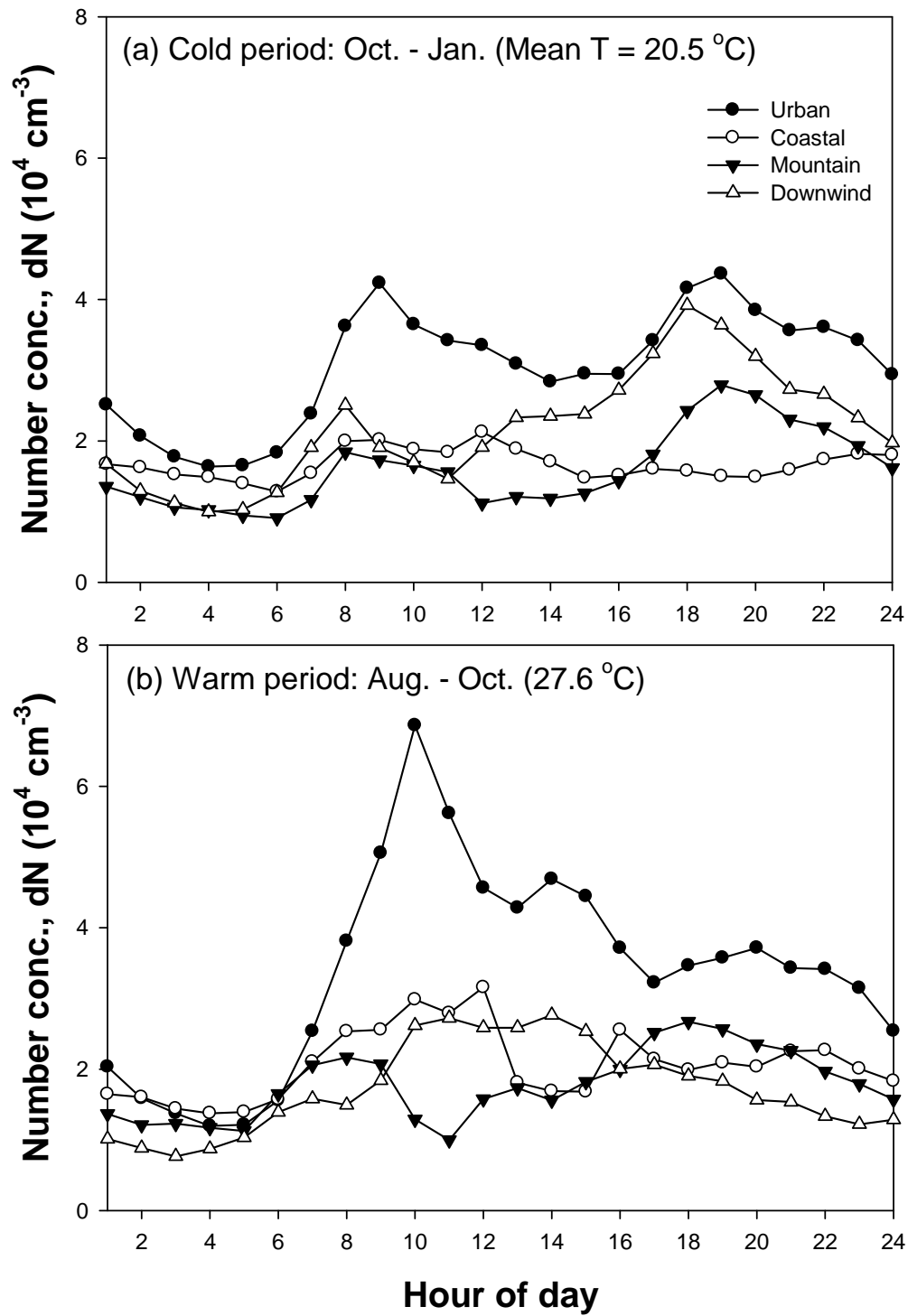
4 **Fig. 4.** Statistical presentations and tri-modal lognormal fits of the particle number size

5 distributions at the (a) urban, (b) coastal, (c) mountain and (d) downwind site during the

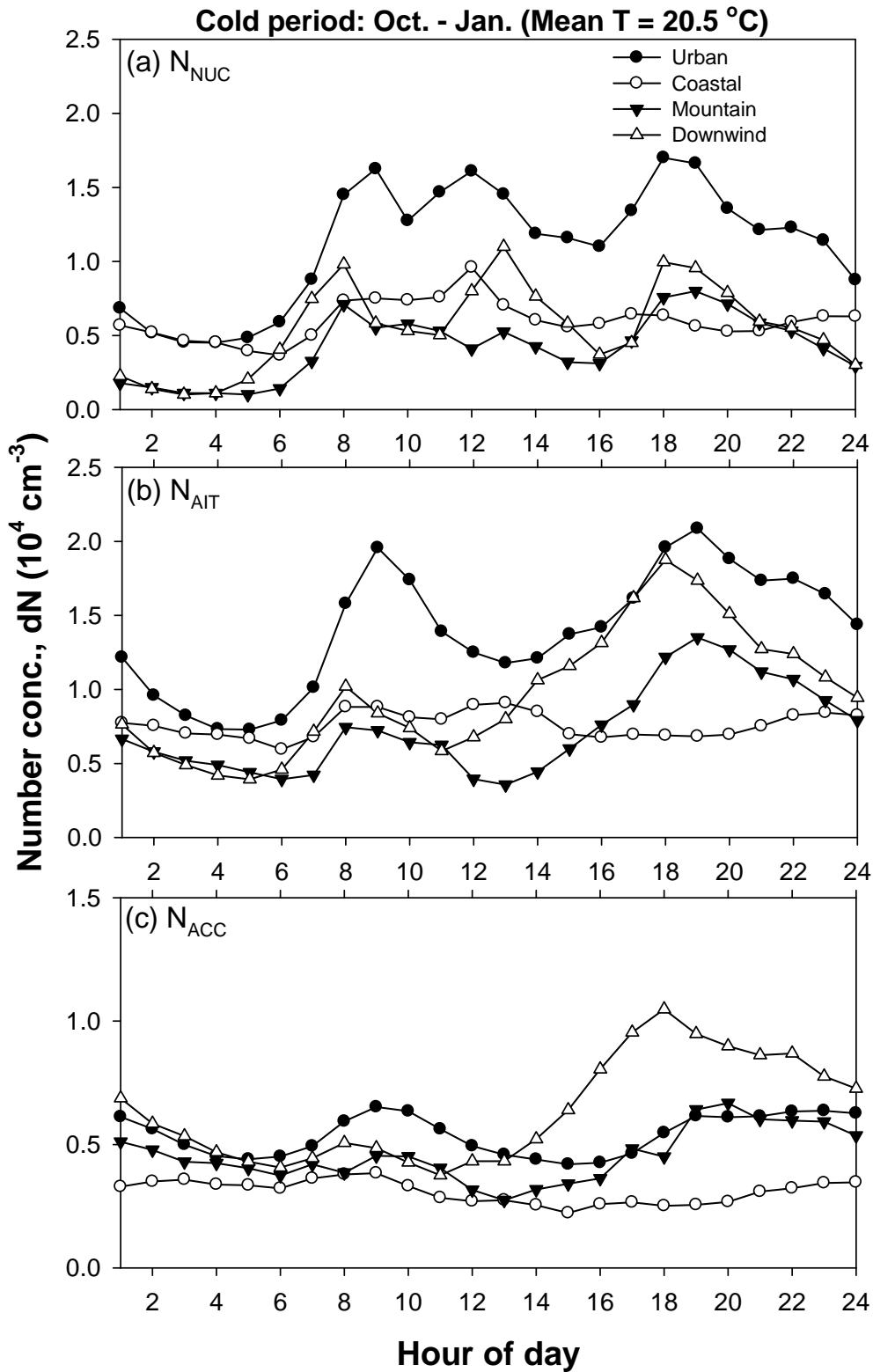
6 warm period.

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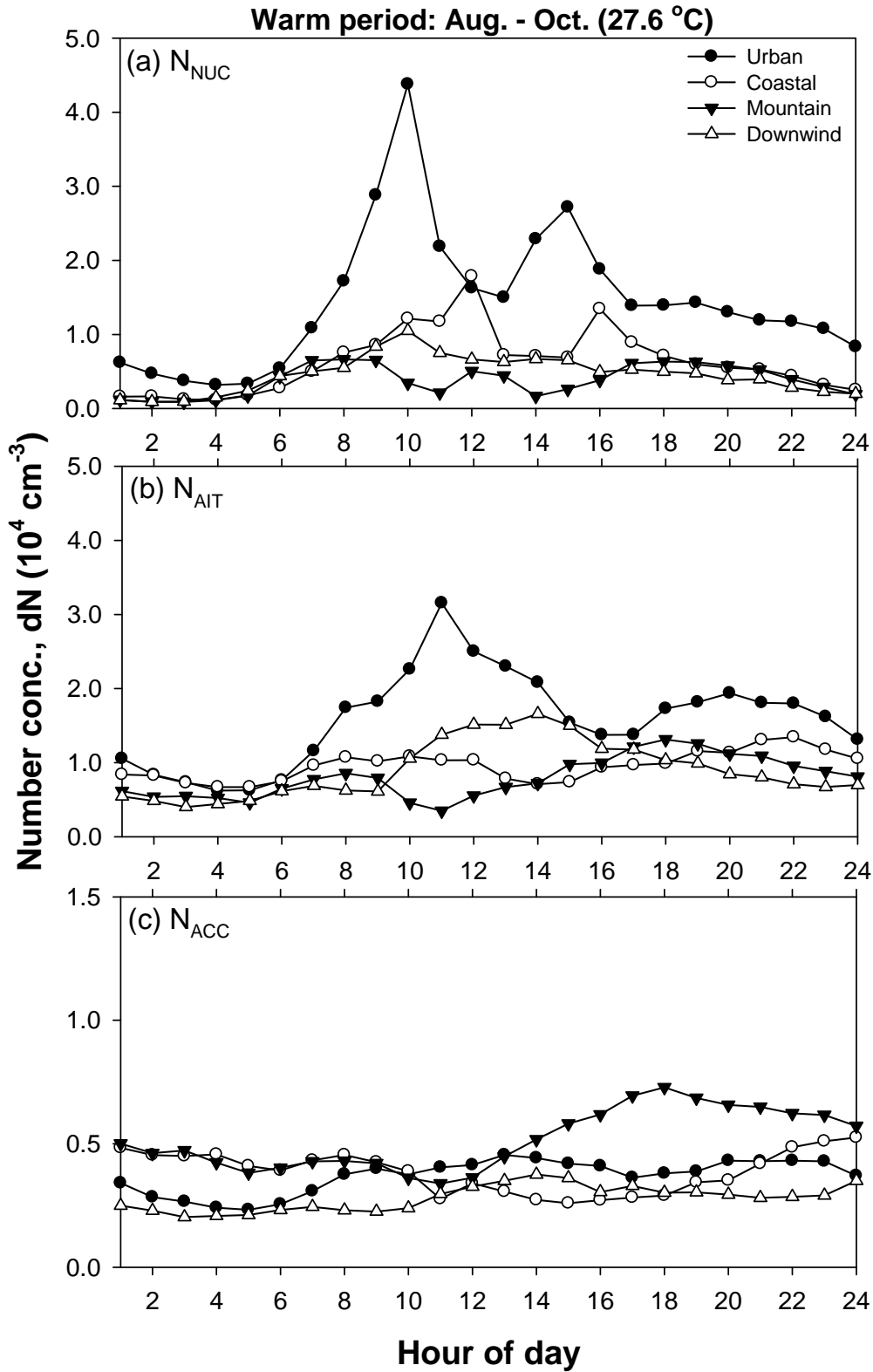
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 2 **Fig. 5.** The diurnal variability of the hourly submicrometer particle number
 3 concentrations (N_{TOT}) at each site during the (a) cold and (b) warm period.



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 2 **Fig. 6.** The diurnal variability of the hourly (a) N_{NUC} , (b) N_{AIT} and (c) N_{ACC} at each site
 3 during the cold period.



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 2 **Fig. 7.** The diurnal variability of the hourly (a) N_{NUC} , (b) N_{AIT} and (c) N_{ACC} at each site
 3 during the warm period.

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SUPPLEMENTAL MATERIAL

**Spatiotemporal variability of submicrometer particle number
size distributions in an air quality management district**

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2 **TABLE AND FIGURE CAPTIONS**

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4 **Table S1.** Averages of daily mean total submicrometer number concentrations (N_{TOT} , 10^4
5 cm^{-3}) on weekdays and weekends during the two study periods.

6

7 **Fig. S1.** The site-specific diurnal variability of wind speed during the (a) cold and (b)
8 warm period.

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10 **Fig. S2a.** The site-specific diurnal variability of air pollutants during the cold period.

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12 **Fig. S2b.** The site-specific diurnal variability of air pollutants during the warm period.

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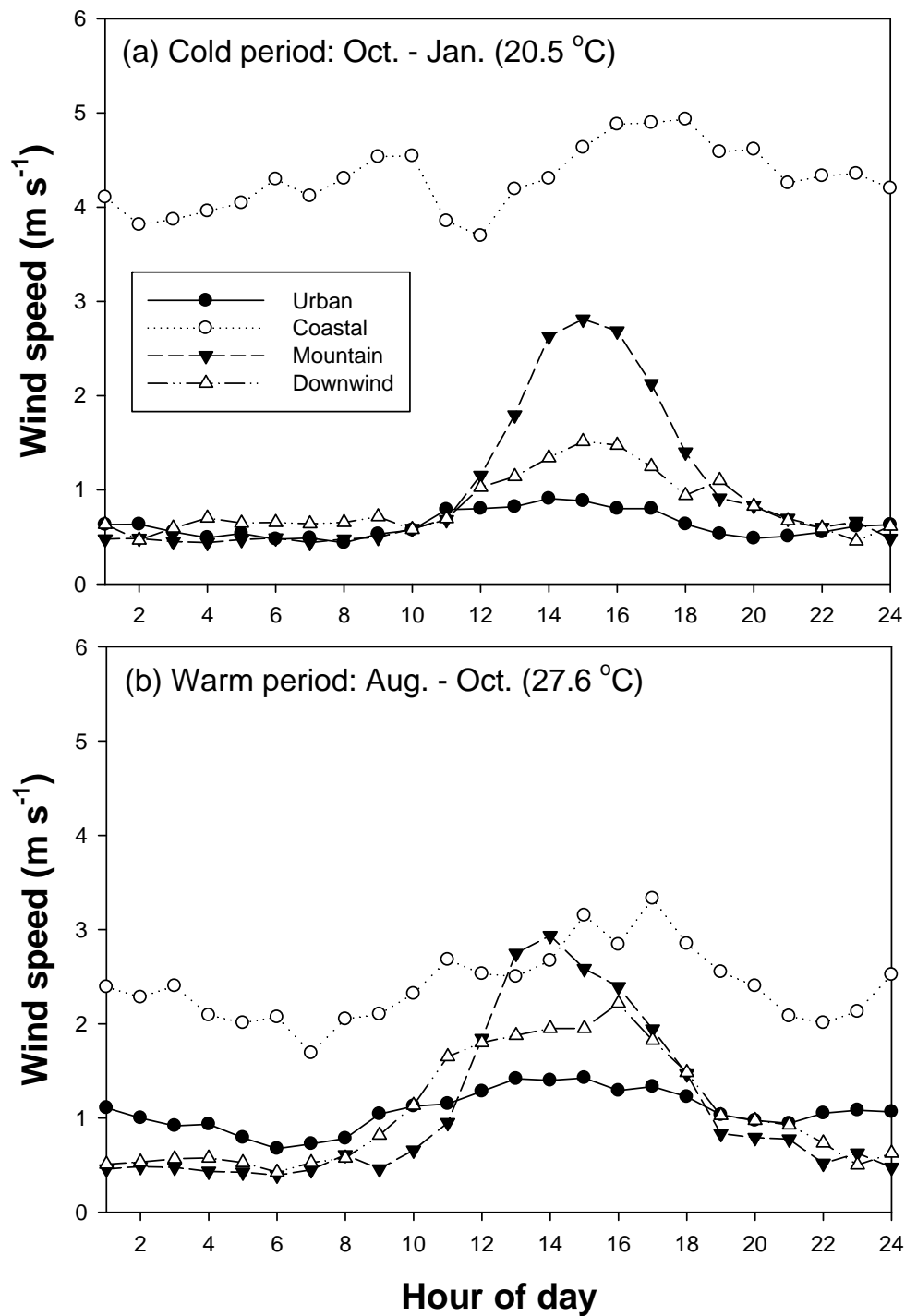
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Table S1. Summary statistics of daily mean total submicrometer number concentrations (N_{TOT} , 10^4 cm^{-3}) on weekdays and weekends during the two study periods.

Period	Weekday N_{TOT}			Weekend N_{TOT}		
	n	Average	SD	n	Average	SD
Cold	67	2.3	0.9	24	2.3	1.2
Warm	36	2.3	0.8	10	1.9	0.8
Overall	103	2.3	0.9	34	2.2	1.1

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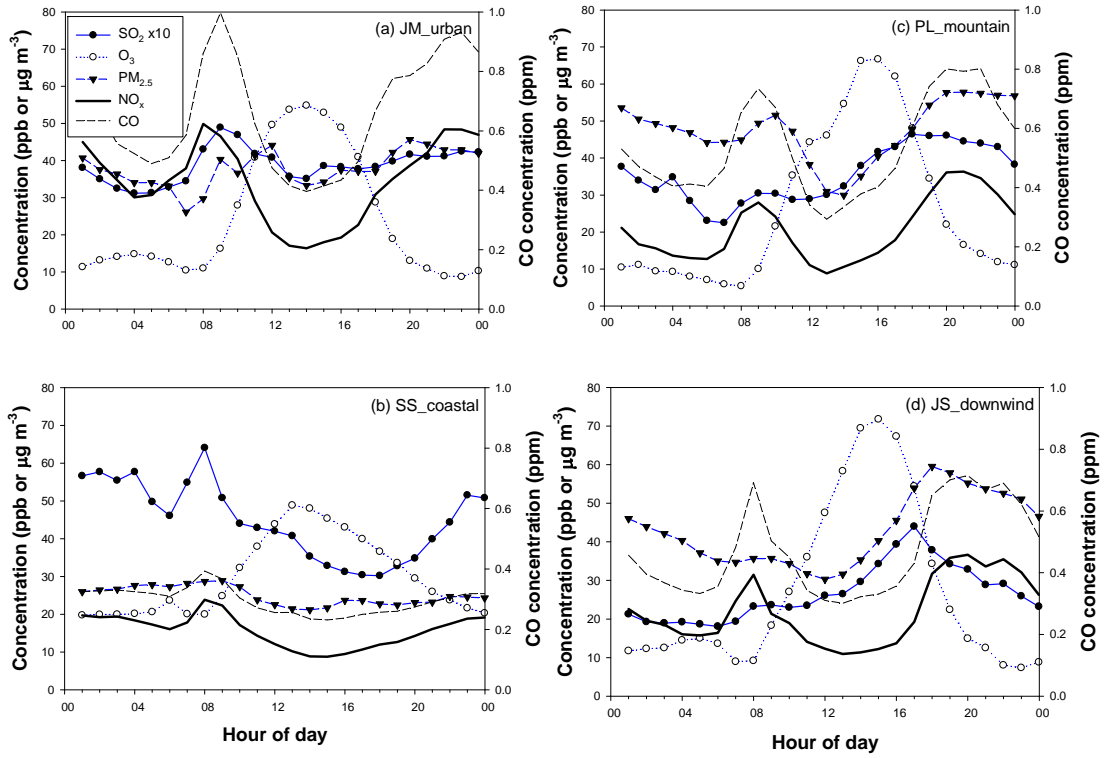


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Fig. S1. The site-specific diurnal variability of wind speed during the (a) cold and (b) warm period.

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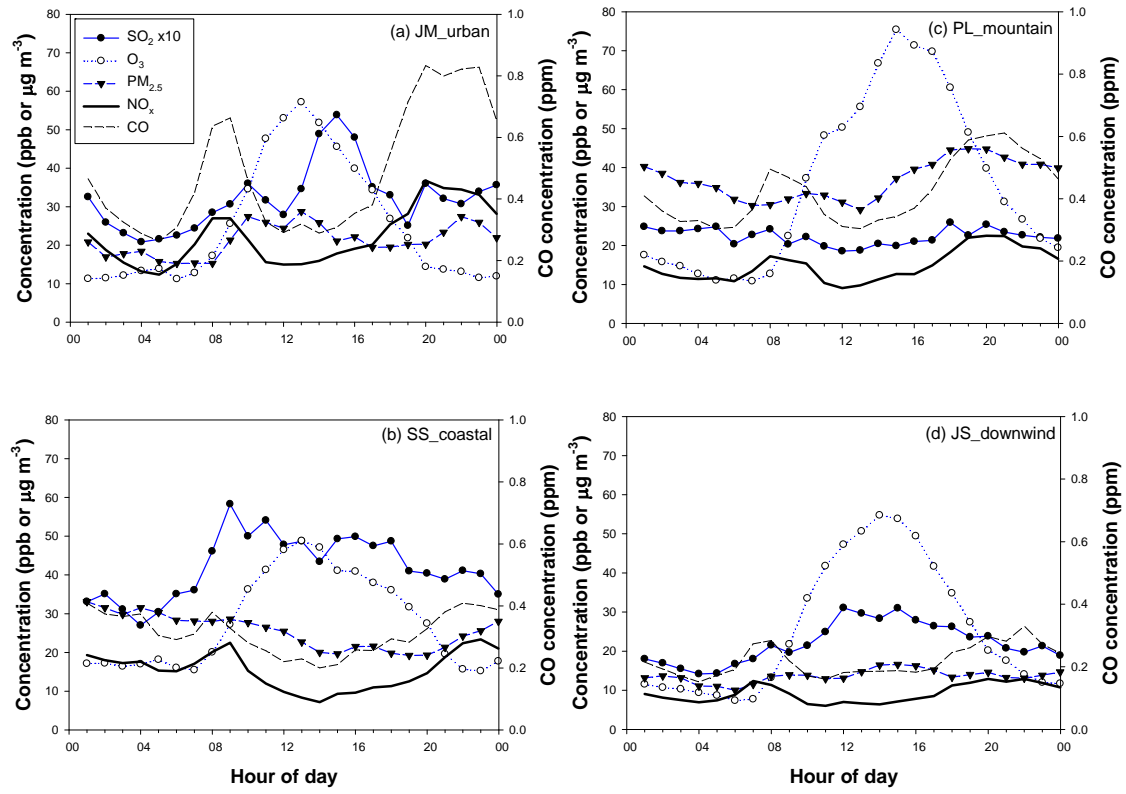
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Fig. S2a. The site-specific diurnal variability of air pollutants during the cold period.

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Fig. S2b. The site-specific diurnal variability of air pollutants during the warm period.